

#### **Abstract**

Continuing development of coal bed gas (CBG) resources has prompted concerns from local government and the public about the potential impacts to ground water in the Raton Basin. There is a concern of possible de-watering of shallow aquifers due to water being pumped from CBG wells. Some people see the CBG waste water as a valuable resource that is being lost in an arid basin. There is also a safety concern that continued development of CBG may cause gas seepage from coal seams and coal sub-crops. This presentation summarizes methods useful to address such concerns.

A total of 2,749 linear miles were surveyed for hydrocarbon seeps using a truck-mounted infrared spectroscopy detection device. Sixty seven (67) separate seeps were found, and these were confined to an area within the Raton Formation coal outcrop and shallow subcrop. Water wells containing measurable amounts of dissolved methane are geographically distributed in a pattern similar to that observed for methane seeps. Of 240 water well samples analyzed, 47% contain measurable amounts of dissolved methane, 27% contain more than 2 mg/L of methane, and 8% contain more than 10 mg/L of methane. The source of methane in these wells can be determined by analyzing its stable carbon and deuterium isotope values. The natural range of stable isotopic values found is very large and allows the methane in groundwater and shallow seeps to be differentiated from produced CBG methane.

Four hundred and sixty five (465) active and abandoned coal mines were identified, three hundred and twenty eight (328) were digitized, and the location of one thousand one hundred and forty one (1141) core holes were digitized.

The quality of surface water, domestic well water, and springs is highly variable throughout the basin and composed of mixed water types. These types include calcium-magnesium bicarbonate (dissolved limestone), calcium and sodium sulfate (dissolved gypsum and thenardite), and minor amounts of sodium bicarbonate (dissolved baking soda). Water quality and water well yield is strongly affected by naturally-occurring coliform, fecal coliform, sulfate-reducing, and iron-related bacteria. Analysis of the distribution of dissolved sulfate, a characteristic of shallow aquifers, indicates that mixing between waters from shallow aquifers occurs within a vertical interval of 400 feet. This may have implications for the depths needed to set surface casing.

Produced waters contain significantly higher levels of total dissolved solids and are composed principally of two mixed water types: sodium bicarbonate (dissolved baking soda) and sodium chloride (dissolved table salt). Analysis of the spatial and temporal distribution of water types among producing wells in the basin suggests that lateral communication between aquifers is not likely to exceed a distance of 1.5 miles. This result offers useful information needed to address the need for groundwater monitoring along outcrop belts, particularly in areas where wells produce abundant water. A few CBG wells were found to produce water containing dissolved sulfate. Such wells need to be analyzed further to determine whether they may be tapping shallow water sources. The stable isotopic composition of carbon in dissolved inorganic carbon is also a useful indicator for detecting the presence of shallow groundwater in producing wells.

# THIS STUDY ADDRESSES SIX WATER & GAS-RELATED ENVIRONMENTAL ISSUES ASSOCIATED WITH COALBED GAS DEVELOPMENT

- METHANE SEEPS
- DISSOLVED METHANE IN GROUNDWATER
- NATURAL GROUNDWATER QUALITY
- POTENTIAL IMPACT TO SHALLOW AQUIFERS
  - WATER QUALITY
  - GROUNDWATER LEVELS
- SCREENING METHODS, FORENSIC TOOLS, & QUALITY CONTROLS AND PROTOCOLS

#### **MULTIPLE DATA SETS USED IN STUDY**

### **DATA**

- COGCC Survey
  - SeaCrest
  - SeaCrestESN Rocky Mtn.
  - Isotech
- KLT
- PETROGLYPGH APPLIED ECO.

### COAL MINE INVENTORY • AMOCO

COGCC Survey

#### PRODUCED WATER SEEP & GROUNDWATER **DATA**

- COGCC Survey
  - Seacrest
  - ESN Rocky Mtn.
    - Apogee Scientific
- USGS WSP 2288
- CEDAR RIDGE
- EVERGREEN

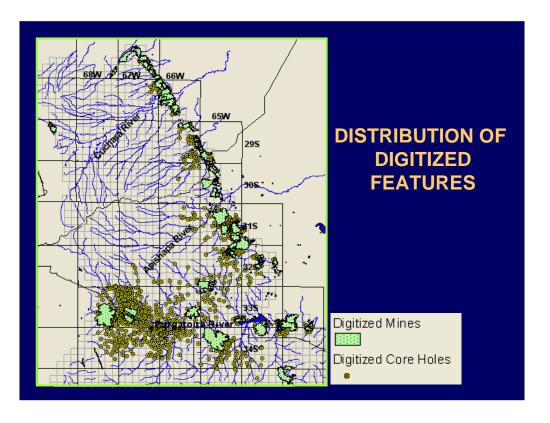
### OBJECTIVES OF COGCC BASELINE STUDY

- COLLECT, COMPILE, AND ANALYZE A VARIETY OF ENVIRONMENTAL BASELINE INFORMATION
- MAKE DATA AVAILABLE TO STAFF, INDUSTRY & PUBLIC
- ADDRESS POTENTIAL ENVIROMENTAL RISK TO SURFACE AND SHALLOW GROUNDWATER RESOURCES WHICH MAY IMPACT STAKEHOLDERS
- EXAMINE DOMESTIC AND PRODUCED WATER QUALITY
- ESTABLISH METHODS TO IDENTIFY, QUANTIFY, AND RANK RISK
- ASSESS RISK OF COALBED GAS OPERATIONS AT THE SCALE OF THE BASIN
- RECOMMEND PROTOCOLS FOR FUTURE MONITORING, SAMPLING, AND ANALYSIS

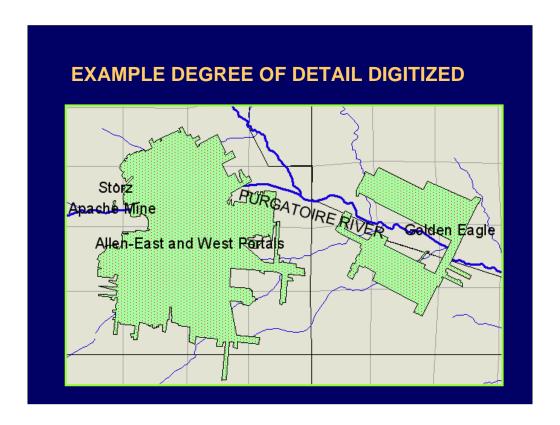
This baseline study emphasizes tools and techniques that are useful for addressing complaints and for evaluating the potential environmental impact of coalbed gas production operations on groundwater resources. No attempt is made to either define risk, rank risk, or to list and rank the potential consequences should any risks be realized. Risk, as defined here, is neutral term used for evaluating the uncertainty of an outcome having an undesired effect.

### COAL MINE INVENTORY DIGITIZED: Lewicki & Associates

- •465 MINES IDENTIFIED
- •328 MINE BOUNDARIES DIGITIZED
- •MINE PORTALS AND OTHER FEATURES
- THAT MAY EMIT METHANE DIGITIZED
- **•LOCATION OF 1411 CORE HOLES DIGITIZED**



The core hole data base marks locations only. Lithologic data have not been digitized.



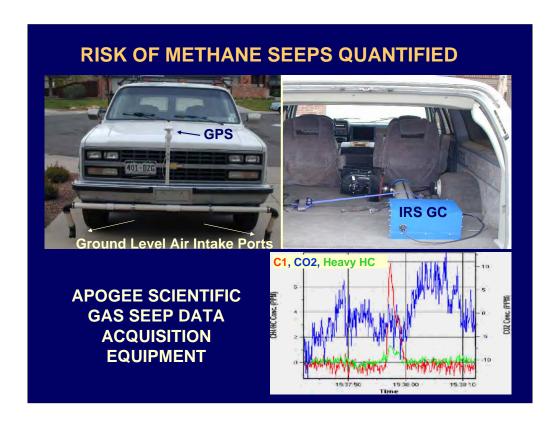
Detailed mapping of mines helps as follows:

- •Provides guide for placement of water wells and producing wells
- •Identifies mine portals which may be source of methane mine emissions.

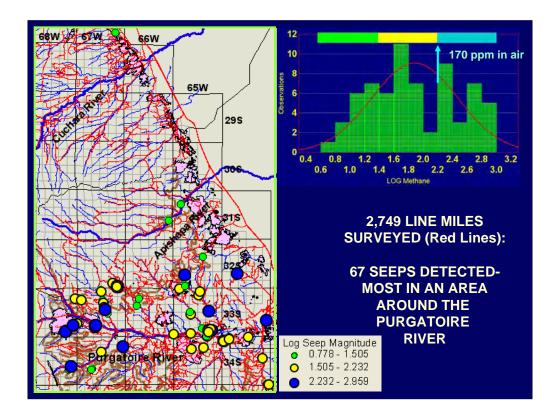
### METHANE SEEPS & DISSOLVED METHANE DOCUMENTING BASELINE CONDITIONS

**APOGEE SCIENTIFIC** 

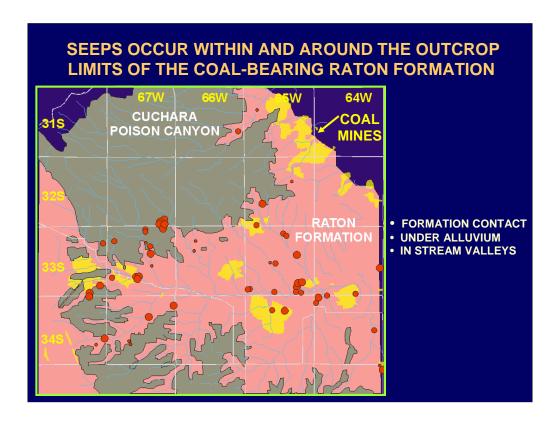
- METHOD USED FOR DETECTING SEEPS
- SEEP DISTRIBUTION AND MAGNITUDE
- RECOMMENDATIONS FOR ADDITIONAL SURVEYS
- DISTRIBUTION OF DISSOLVED METHANE
- METHOD USED TO IDENTIFY METHANE SOURCE



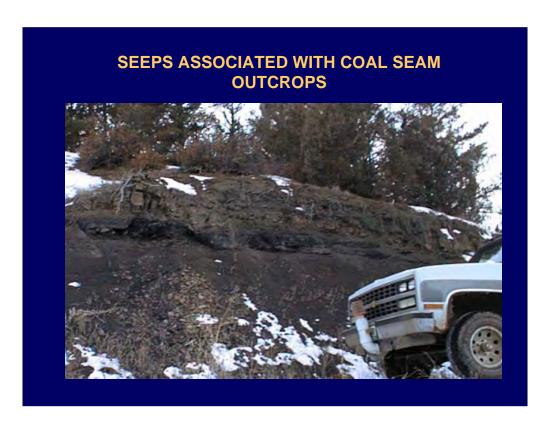
The gas detection system is based on infrared spectroscopy which differentiates methane, carbon dioxide, and heavy hydrocarbons. Air scoops, located at the front of the vehicle, continuously sample air which is analyzed at a maximum rate of 100 Hz. Sensitivity to hydrocarbons of the IRS device is on the order of a few ppm by volume in air. Proximity to seeps is estimated by examining peak heights and widths. The sharper the peak, the closer the seep. Once detected, seeps can be pinpointed by driving in circles.



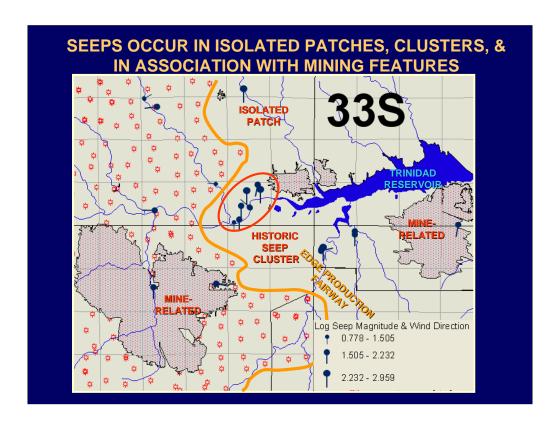
Methane seeps detected in this baseline study are clustered within the Purgatoire River drainage basin. Red lines indicate the 2749 mile route taken for the survey. The northernmost seep occurs at the mouth of a mine. A histogram of the logarithm of seep concentrations (in ppm) shows there are two seep populations: one less than approximately 169 ppm; the other greater than 169 ppm. For reference, the lower explosive limit (LEL) for methane in air is 5% by volume, or 50,000 ppm (log value 4.69). Safety levels of concern are usually set at 20% of the LEL, or 10,000 ppm (log value 4)



Geologic mapping shows that seeps occur exclusively within the area defined by the outcrop and shallow subcrop of the coal-bearing Raton Formation. 67 seep locations were detected. This is a small number considering the relatively vast areal extent of the Raton Formation outcrop belt. Seeps emanating from the coal-bearing subcrop are buried at depths less than 50 feet below either alluvium, or the Poison Canyon Formation. There is no apparent correlation between seep locations, coalbed gas fields, mapped dikes, or mapped faults. The majority of natural seeps occur within and along stream valleys and along road cuts where coal seams are exposed.



Example of a seep emanating from exposed outcrops.



These tadpole plots show relative seep magnitudes as well as wind direction monitored at the time of the survey. Historic seeps occur outside of the updip limit of coalbed gas production activities. Coalbed gas wells are marked in red. No attempt was made to identify the source of seeps occurring within the production trend. Pinpointing seeps would be best accomplished by walking around a detected gas source using a hand-held flame ionization detection device.



Approximately 1/3 of the seeps detected appear to be emanating from mine-related features.

### SUMMARY OF SEEP CHARACTERISTICS

### ANTHROPOGENIC ORIGIN

- Mine-related features
- Water wells
- Pipeline right-of-ways
- Other utility Infrastructures
- Gas production
- Historic

### NATURAL ORIGIN

- Within outcrop trace of coal-bearing formation
- Formation contacts
- Dominant in NW & NEtrending linear stream valley traces
- Outcrop buried under shallow alluvial fill
- Historic

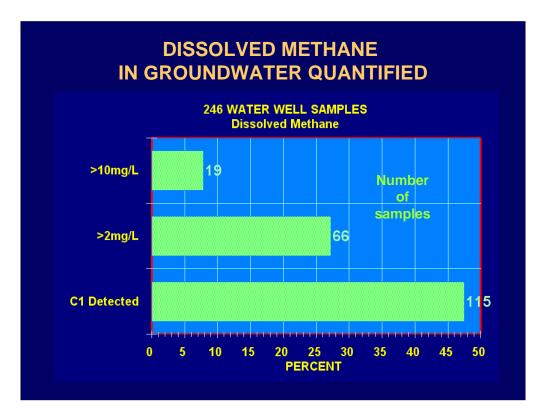
Based on the results of the survey and interviews with some land owners, only one seep cluster detected may possibly be associated with commercial gas production activities. Further investigation is required to determine the source of methane in this area.

#### IRS IS A GOOD SCREENING TOOL

**COVER A LOT OF AREA IN SHORT TIME** 

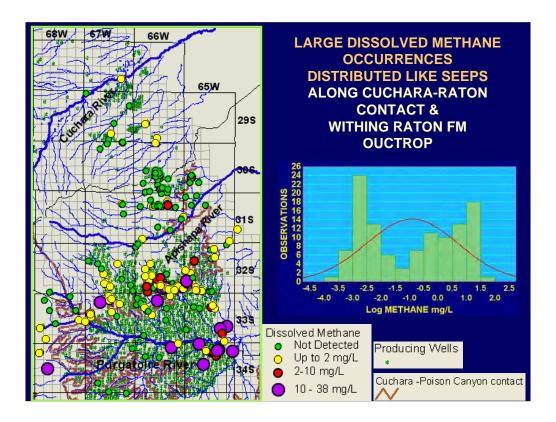
#### RECOMMENDATIONS

- EXTEND IRS SURVEY TO OUTCROP BELT NEAR RECHARGE AREA OF SOUTH FORK PURGATOIRE
- FOLLOW UP SCREENING TOOL WITH WALKING SURVEY USING PORTABLE FID
- LOCATE AND DESCRIBE CHARACTERISTICS OF SPECIFIC SITES: CLASSIFY SOURCE
- SAMPLE GAS FROM LARGER SEEPS FOR LABORATORY ANALYSIS & COMPARE TO PRODUCTION WHERE NEEDED
- DETERMINE NEED FOR LONG TERM SOIL GAS FLUX MONITORING

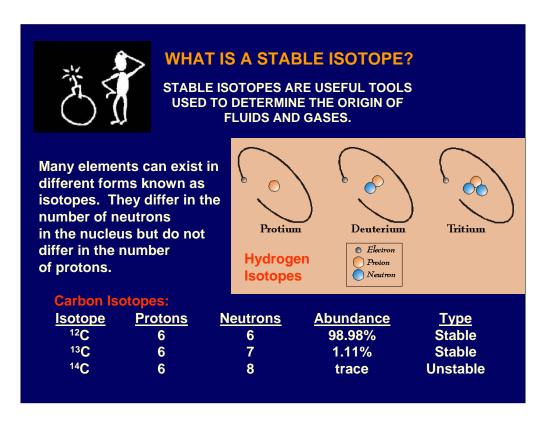


Nearly half of all water wells tested for methane contain detectable concentrations. This result is normal for foreland basins along the Rocky Mountains. 27% of the wells tested contain concentrations exceeding 2 mg/L. Such concentrations define an arbitrary threshold where methane could exsolve in enclosed spaces and accumulate to explosive levels. 7% of the wells tested contain concentrations exceeding 10 mg/L. For comparison, the saturated dissolved methane concentration at atmospheric pressures typical for the high elevations in this basin is approximately 24 mg/L. Wells containing more than 10 mg/L of dissolved methane should be treated to safely remove exsolving methane.

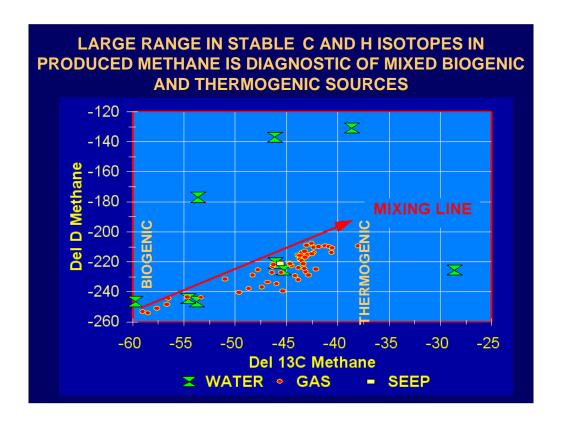
Wells should be screened for methane by filling a clean 5 gallon bucket with water and testing the air above the water with a hand-held flame ionization detector. If methane is detected, a water sample can be collected by filling a 40 ml. VOA vial under water, and ensuring that there are no bubbles present when the vial is sealed. Such samples are then sent to a laboratory for analyses defined by the USGS Kampbell protocol. Samples collected for chromatographic and stable isotopic analyses are collected by filling a 1L glass amber bottle under water. All efforts should be taken to insure that the bottle does not contain any air bubbles when sealed with a Teflon-coated rubber seal cap. If the water effervesces, then a sample should be collected in the 1L amber bottle by filling the bottle with water under water, using an inverted funnel to allow gas to displace water, and creating a headspace that fills at least ¼ of the bottle volume. Samples are sent to the laboratory for analysis after sealing the Teflon-lined cap with tape and shipping bottles chilled and safely packed in an ice-filled chest.



Like methane seeps, the distribution of dissolved methane concentrations is bimodal. The distribution of larger dissolved methane concentrations is confined to the area bounded by the Raton and Cuchara-Poison Canyon contact. Statistical analysis of the data indicates a mean value of 3 mg/L and a standard deviation of 5 mg/L for samples in the second population of data. Such values could be used to set regulatory requirements for isotopic analysis and monitoring similar to requirements now in place in the San Juan Basin. For example, water well samples found to contain more than 3 mg/L of dissolved methane should be sampled for stable isotopic analysis to determine its source. Changes in dissolved methane concentration of more than 5 mg/L could be used to define requirements for regular water well monitoring if the methane is found to be isotopically similar to that being produced in nearby coalbed gas wells.

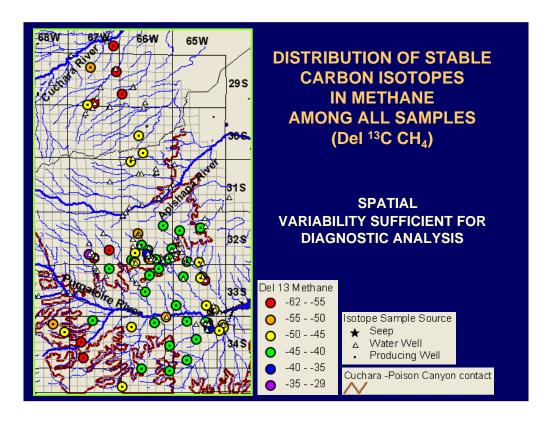


The stable carbon and deuterium isotopes of methane (CH<sub>4</sub>) are used in the oil and gas industry as a diagnostic tool to determine methane origin. Methane originating from the burial of organic sedimentary matter at high temperatures and pressures is defined as thermogenic. Methane originating from bacterial fermentation or from the bacterially-mediated reduction of carbon dioxide is defined as biogenic. Biogenic methane is a common constituent of groundwaters around the world. Approximately 20% of the world's commercial natural gas reserves is biogenic.

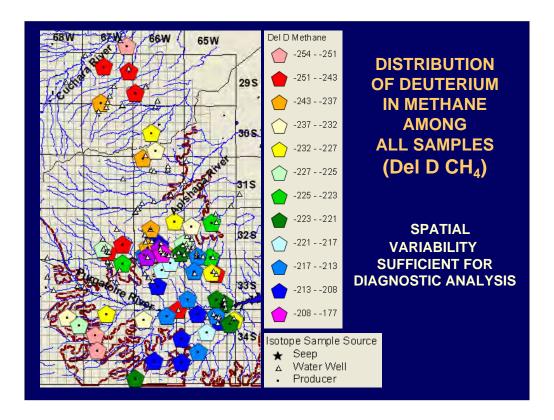


The isotopic pattern evident among samples collected from coalbed gas wells in the Raton Basin is one that is characteristic of mixed biogenic and thermogenic sources. The gas composition of samples with isotopic signatures characteristic of thermogenic methane generally contains higher hydrocarbons such as ethane, propane, butane, and in some cases, pentane. The gas composition of samples with isotopic signatures characteristic of biogenic methane generally contain only methane and minor amounts of ethane. Stable isotopic analysis of carbon in produced  $CO_2$  yields very positive values, indicating that bacteria have converted  $CO_2$  into methane. There is little apparent correlation between the isotopic composition of methane and the vitrinite reflectance in coals. This indicates that the gas composition has been altered by bacteria which may have been introduced along basin margins during uplift and fresh water intrusion into the basin. In general, wells producing gas with a thermogenic signature produce less water and more gas than wells producing gas with a biogenic signature. Water from wells producing thermogenic gas also tends to be more saline and enriched in chlorides.

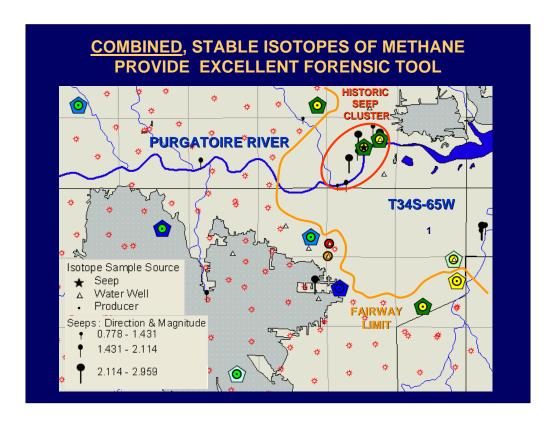
There are only a few complete stable isotopic measurements available for methane in domestic water wells, springs, and seeps. Those available are plotted above. In some cases, the isotopic composition is similar to that observed among producing wells. In other cases, the signatures are very different. Because of the large range in isotopic values, stable isotopic analysis can be a very useful forensic tool to determine the origin of methane in domestic water wells, springs, and seeps.



In most places, the stable isotopic composition of carbon in methane from shallow water wells is distinctly different from that found in nearby producing wells.



The spatial distribution of deuterium values in methane from produced wells is more variable than the distribution of stable carbon values. This probably reflects a more heterogeneous distribution of the deuterium in water from which the methane was derived.

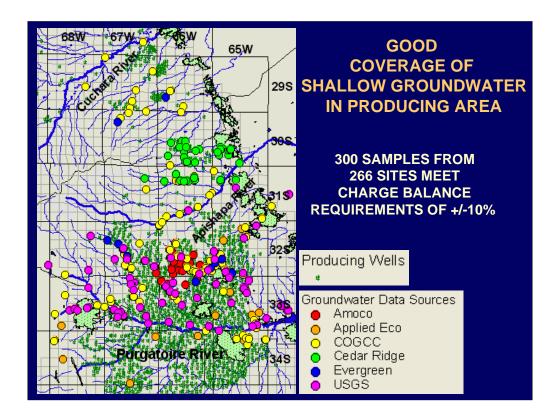


Combined, the stable carbon and deuterium isotopes of methane can be used to determine the origin of methane in seeps and domestic water wells. For example, the historic seeps at the mouth of the Trinidad Reservoir and dissolved methane in a nearby domestic water well have a composition that is nearly identical to the composition of gas from the producing well in section 28, T33S-R65W. This is an area that could be monitored over the long term to determine whether the intensity of seeps is increasing as a result of down-dip production. In the SW corner of section 2 in T34S-R65W, the stable isotopic signature of carbon in domestic water wells is very different from that in a nearby producing well in section 3. Methane in the domestic water wells here probably originates from shallow biogenic sources.

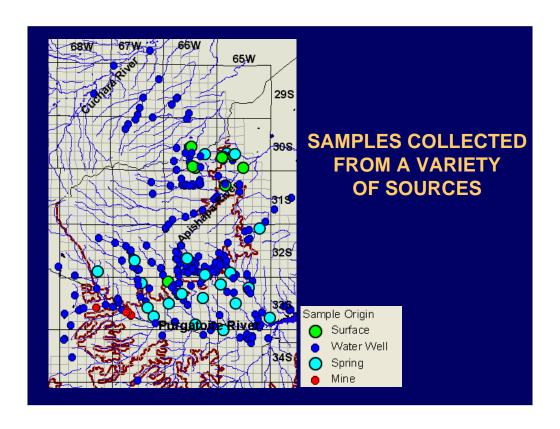
### **SUMMARY DISSOLVED METHANE**

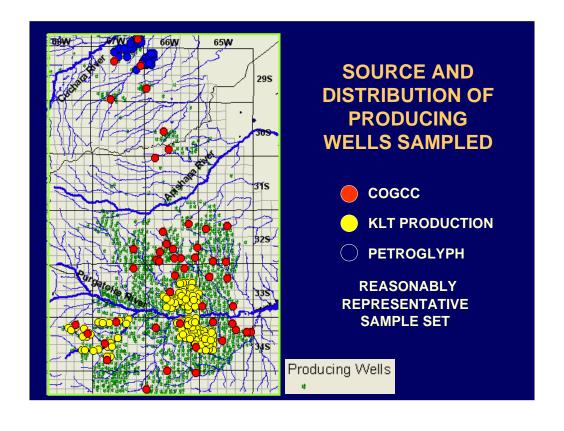
- DISSOLVED METHANE OCCURS WITHIN OUTCROP BOUNDARY OF RATON FORMATION
- REGULAR MEASURMENT OF DISSOLVED METHANE IN WATER WELLS RECOMMENDED
- STABLE ISOTOPES USEFUL FOR DETERMINING METHANE SOURCE AND FOR COMPARISON WITH GAS FROM PRODUCING WELLS

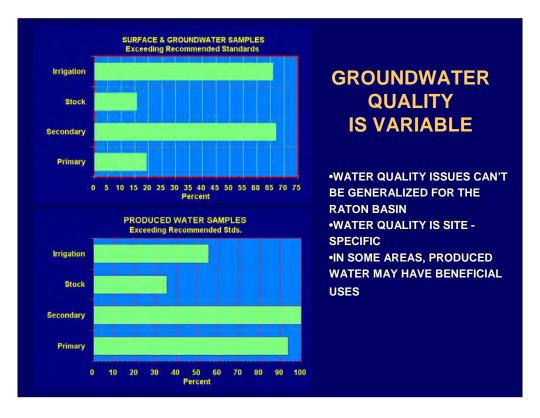




The quality of water analyses has been checked by comparing the sum of the concentrations of negatively charged dissolved anions against the sum of the concentrations of positively charged dissolved cations. Theoretically, the sum of positive and negative charges should be zero. However, due to the different analytical methods used to analyze positive ions and negative ions, it is customary to accept analyses that are within a charge balance of +/- 5%. Because of differences in analytical methods applied to obtain values used for this study, the use of different laboratories, and systematic sampling errors found in the data analyzed in this study, we chose to accept analyses that had a charge balance of +/- 10%. The largest source of errors detected in this study originate from the titration methods used to measure alkalinity. Ideally, samples should be titrated in the field because the loss of dissolved CO<sub>2</sub> during transport and protracted sample storage reduces *in-situ* values by as much as 10%. If samples are not titrated in the field, they should be titrated within one or two days after arriving at the analytical laboratory.



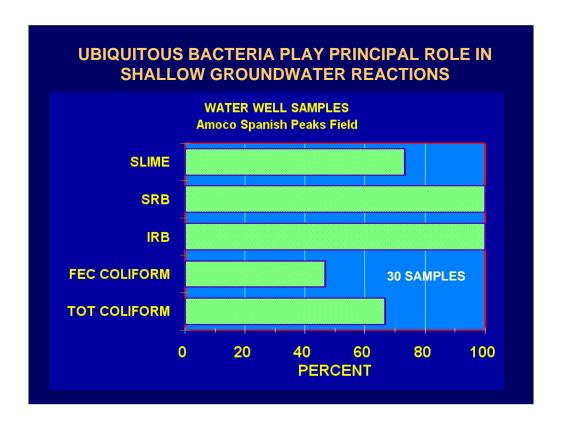




Water quality from domestic water wells is compared to water quality of produced water samples in this illustration. The majority of domestic water well samples exceed recommended secondary drinking and irrigation water standards. The majority of produced water samples exceed recommended primary and secondary drinking water standards. The primary cause for the failure to meet primary and secondary drinking water standards is high salinity and a predominance of sodium-rich waters.

Domestic water well samples did not contain unacceptable concentrations of the following dissolved RCRA metals: selenium, mercury, copper, chromium, cadmium, barium, arsenic, and antimony. Less than 1% of the samples analyzed contained lead levels in excess of recommended drinking water standards. 13% contained unacceptably high sodium values, and approximately 6% of the samples contained unacceptable concentrations of fluoride and nitrates. 55% of samples analyzed exceeded irrigation standards for molybdenum.

Barium was the only RCRA metal detected in produced water samples with 7% of wells tested exceeding recommended drinking water standards for this element. 45% of wells tested contained fluoride concentrations in excess of drinking water standards, and 95% contained sodium concentrations in excess of drinking water standards. The percentage of wells exceeding irrigation standards illustrated here is based on high values of TDS and pH only. Average sodium adsorption ratios (SAR) for produced waters is 70, making the majority of produced waters unfit for irrigation purposes.



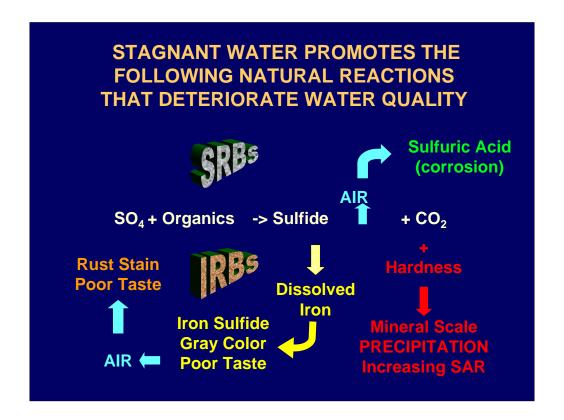
The data shown above are derived from a limited set of Biologic Activity Reaction Tests (BART). Because BART samples are easy to collect and analyze, they should become a standard part of domestic water well sampling protocols. Water wells in most western basins contain bacterial consortia that can significantly alter water quality and water well yield.

#### **BACTERIAL CONSORTIA FOUL WELLS**

- CONSUME OXYGEN AND TURN WATER STAGNANT
- BIOSLIME CLOGS PERFORATIONS
- MINERAL SCALING CLOGS PERFORATIONS

**RESULT:** 





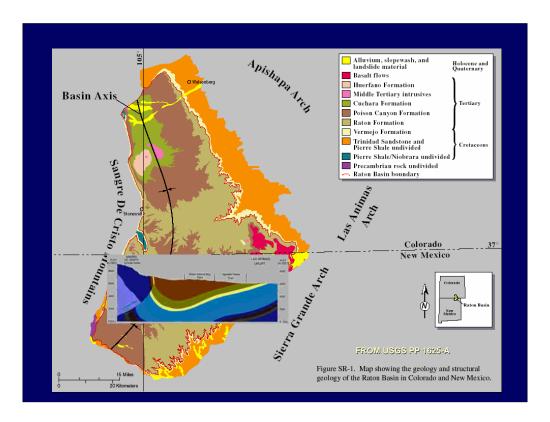
The principal cause of poor domestic water quality and low water well yields in western basins is bacterial fouling and mechanical problems. There are established methods available to determine whether problems with water quality and well yield are related to bacteria. Analytical methods which can be used to sample wells in the field include BART (Biologic Activity Reaction Tests), analysis of fixed gases, and analyses of chemical composition. Results are used to determine the relative impact of scaling and bio-fouling. Chemical analysis is used to determine the scaling or corrosion potential of water (e.g. Langelier Saturation Index and Ryznar Stability Index). Such tests should be conducted and analyzed routinely before attempting to determine whether changes in water quality and yield are due to coalbed gas or other oil and gas production activities.

## ADDRESSING COMMUNICATION BETWEEN SHALLOW AQUIFERS AND GAS PRODUCING ZONES

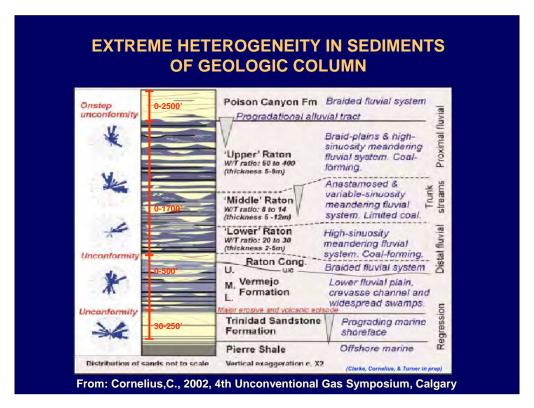
- PUBLIC CONCERN THAT PRODUCING WATER FROM COAL IS IMPACTING AVAILABILITY OF WATER FROM SHALLOW AQUIFERS
- NO REGIONAL, SYSTEMATIC MONITORING OF SHALLOW GROUNDWATER LEVELS ESTABLISHED IN THE BASIN
- NEED TO DEVELOP SCREENING METHODS USING AVAILABLE DATA THAT IDENTIFY AREAS WHICH MAY BE AT RISK OF CROSS-AQUIFER COMMUNICATION

### APPROACH USED TO DEVELOP SCREENING TOOLS

- MAP AREAS WHERE WATER WELLS AND PRODUCING WELL COMPLETIONS ARE IN CLOSE VERTICAL AND LATERAL PROXIMITY
- IDENTIFY DIAGNOSTIC PARAMETERS THAT DIFFERENTIATE SHALLOW GROUNDWATER AND PRODUCED WATER
- MAP KEY PARAMETERS
  - IDENTIFY HOW THEY VARY ACROSS THE BASIN
  - IDENTY HOW THEY CHANGE WITH TIME
- DETERMINE WHICH PARAMETERS IN PRODUCED WATER MAY INDICATE POSSIBLE CONNECTION WITH SHALLOW AQUIFERS



The Raton Basin is a structurally asymmetric synclinal basin, with coal outcrops along the eastern and western margins. Available data indicate that the predominant area of groundwater recharge occurs along the western margin where elevations are highest and precipitation is most abundant. Because the overall topographic slope tilts towards the east, recharging fluids move generally eastward in the subsurface. Along the flowpath, groundwaters discharge into tributary stream valleys. The principal rivers of the area, the Purgatoire, the Apishapa, and the Cuchara, are "gaining" rivers, meaning that flow is enhanced by groundwater baseflow discharge. Thus the chemical composition of water in rivers at any given time results from the mixing of surface runoff and baseflow. At times of low annual precipitation rates, a larger component of the fluid chemistry is due to groundwater discharge; at times of high annual precipitation rates, the fluid chemistry is fresher.

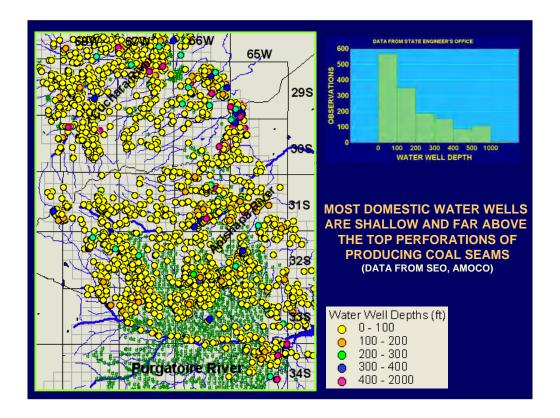


Groundwater flow rates in confined in coal seam aquifers of the Raton and Vermejo Formations are highly variable. Regionally, water pressure measurements indicate that fluid pressure gradients in coal aquifers are below hydrostatic gradients. This is principally due to generally poor vertical and lateral communication between aquifers. The existence of such poor communication is supported by the highly variable aquifer fluid chemistries that will be shown in the next few slides.

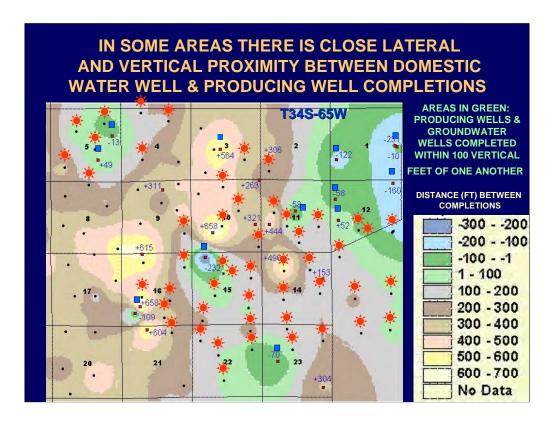
The observed poor lateral and vertical communication between aquifers is caused by a variety of factors. Based on detailed outcrop, core, and well log observations made during the past decade, the vertical and spatial distribution of permeable aquifers is highly irregular. As illustrated above, the distribution of coal and sand layers in the Poison Canyon, Raton, and Vermejo Formations is extremely heterogeneous. This has important ramifications:

- Permeable matrix will not be continuous over large areas;
- •Natural and induced vertical fractures are not likely to propagate large vertical distances because of the predominance of multiple shale barriers;
- We can expect significant regional and local dispersion of fluids flowing through the matrix;
- Aquifers will be strongly compartmented and discontinuous.

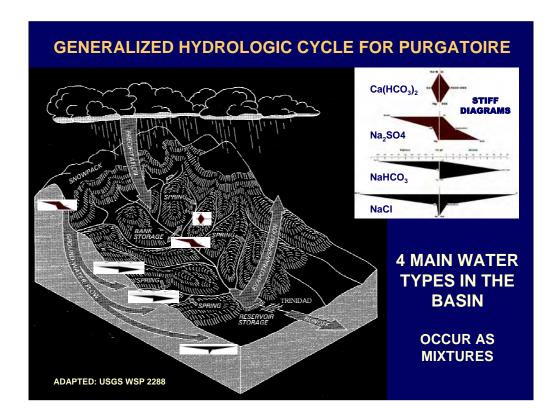
Additional flow barriers are present in the form of intrusive igneous dikes, syn-depositional faults accompanying basin filling and contemporaneous Laramide-age thrusting along the western margin, and small post-depositional faults. Flow rates in sandstones is also very limited by extensive cementation of the available pore space. Coalbed gas production also seals off fluid communication between adjacent aquifers because gas evolving from coal seams significantly reduces the relative permeability of water to negligible rates. This important relative permeability phenomenon is widely overlooked in most assessments of lateral and vertical communication between aquifers. Accordingly, there is little reason to suspect that shallow aquifers will be in natural communication with water produced from deeper horizons in most areas.



The majority of water wells in the basin are drilled to shallow depths of less than 200 feet below the surface. Such depths are far above the upper completion intervals of most producing wells. The extremely heterogeneous geologic framework in the basin makes it very unlikely that there will be vertical communication between shallow aquifer horizons and producing wells, except in areas where the completions in are in close vertical and lateral proximity.



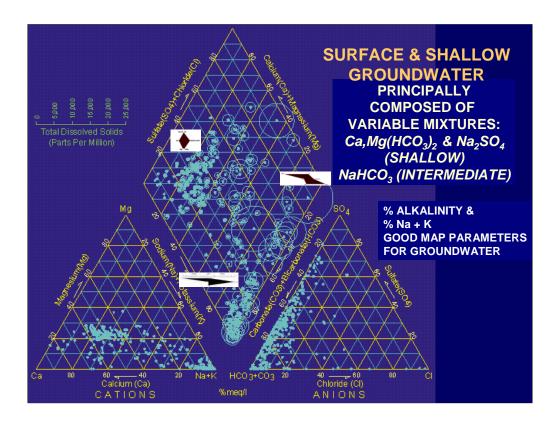
This area of the Raton Basin is one where the depth of water well completions is in relatively close lateral and vertical proximity to the top perforations in producing wells. Water levels in domestic water wells could be systematically monitored here to determine whether there is communication with surface aquifers. Dissolved methane concentrations in domestic water wells could also be monitored to determine if there are systematic increases evident over large areas. Because this area is also one where natural gas seeps have been know to occur and persist long before the advent of coalbed gas operations, it is important to collect and monitor baseline conditions. Only a few number of water level measurements are available in this area. However, the limited data available from areas contoured within the light blue and green range so far do not indicate that there is communication between producing zones and shallow aquifers.



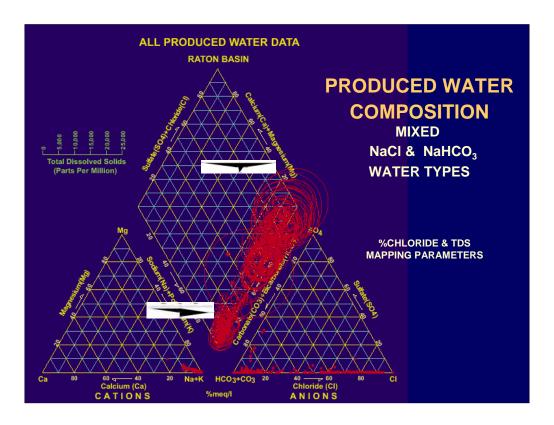
There are four principal water composition types in the basin. These are as follows:

- Dissolved limestone (calcium and magnesium bicarbonate);
- Dissolved thenardite and glauber salt (sodium sulfates) with minor amounts of dissolved gypsum (calcium sulfate);
- Dissolved baking soda (sodium bicarbonate) and;
- Dissolved table salt (sodium chloride).

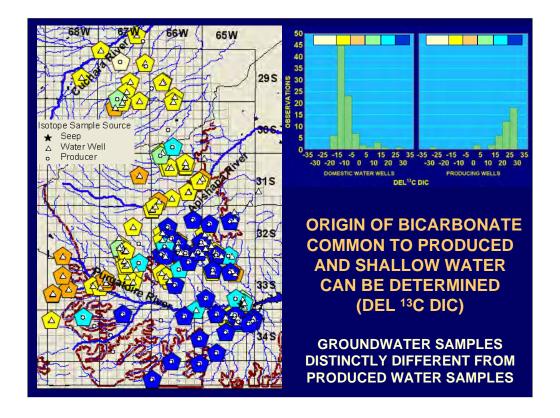
Calcium bicarbonate and sodium sulfate water types form the principal shallow groundwater types, whereas sodium bicarbonate and sodium chloride form the principal intermediate and deep water types. Within the coal-bearing formations, sodium bicarbonate dominates the water composition. Such compositions are typical among coal seams around the world. Sodium chloride forms a significant component of produced waters in the eastern half of the Purgatoire River groundwater flow system. These fluids are most likely present in both coal seams and permeable sandstone aquifers.



A trilinear diagram is used to illustrate the relative composition of the major dissolved constituents in surface and shallow groundwater. Among the positively charged dissolved cations, the distribution of points indicates variable mixtures of two principal endpoint members. End member compositions consist of 80% & 20% calcium and magnesium respectively, and 100% sodium. Among the negatively charged dissolved anions, the distribution of points indicates variable mixtures of two principal endpoint members. End member compositions consist of 80% sulfate and 100% bicarbonate. Calcium-magnesium bicarbonate type fluids contain the lowest amounts of total dissolved solids. Sodium chloride-bearing fluids occur in the most saline domestic water well aquifers, principally located along the eastern half of the Purgatoire River groundwater flow system.

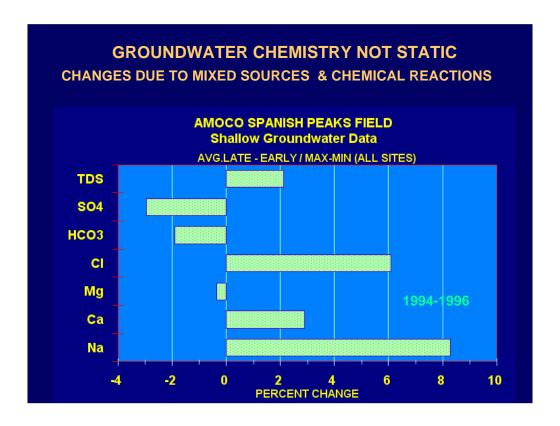


The composition of produced fluids is markedly different from that of surface and shallow groundwaters. Dissolved cations are composed almost entirely of sodium, whereas the dissolved anions are composed of two principal end point members consisting of 100% bicarbonate and 100% chloride.

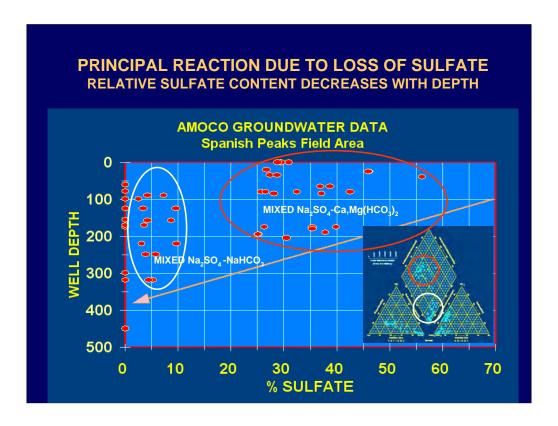


Both shallow groundwaters and produced fluids contain a high relative percentage of dissolved bicarbonate anions. However, bicarbonate can be differentiated on the basis of stable carbon isotopes. The stable isotopic composition of carbon in dissolved inorganic carbon (DIC) among shallow groundwaters is distinctly different from that found in produced fluids. The stable carbon isotope ratios of DIC in shallow groundwater are predominantly in the range between -10 to -15 per mil. This is a characteristic of bicarbonates formed in soils.

The stable carbon isotope ratios of DIC in produced fluids are predominantly in the range of +20 to +30 per mil. These values were derived on the basis of the isotopic composition carbon in the  $CO_2$  analyzed from the gas stream of producing wells. This conversion assumes that gaseous  $CO_2$  has equilibrated with the  $CO_2$  dissolved in water as DIC at an average formation temperature of  $95^{\circ}F$ . Positive values indicate that a significant amount of methane originated from bacteria that convert  $CO_2$  to methane. Bacteria selectively use the lighter carbon isotope in  $CO_2$  for their metabolism (kinetic fractionation), leaving the remaining  $CO_2$  pool enriched in the heavier carbon. Regardless of the origin of such positive values, the large differences observed in the stable carbon isotope values of bicarbonate indicate that this parameter can be used to detect the presence of shallow groundwater in produced fluids.



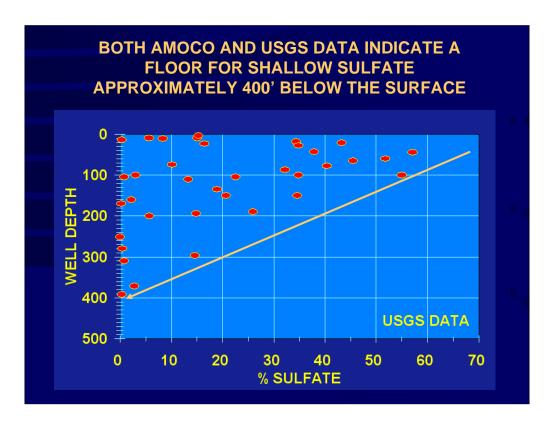
This graph records the variability in concentration of several dissolved species measured in samples collected in 1994 and again in 1996. Paired values of concentrations measured among 25 sample sites were subtracted (value in mg/L 1996 - value in mg/L 1994) and all the values were averaged. This average was divided by the average of the total range of values (maximum - minimum in mg/L) observed in 1994 and 1996 for each dissolved species. One sample data point was omitted because its very high values significantly skewed the averages. Results show a fairly large range in variability at each site relative to the range in values measured among all sites. Such variability is due to variable mixing rates between groundwaters of differing composition. The differences observed are partially a consequence of changes in climatic conditions. At times of low precipitation rates, the relative contribution of discharge from groundwater baseflow increases; at times of high precipitation rates, the relative contribution of discharge from groundwater baseflow decreases. Observed changes are also due to chemical reactions mediated by bacteria. Baseline temporal analyses show that there are natural variations in water quality which must be considered before attempting to determine the potential impact of coalbed gas operations on domestic water well aquifers.



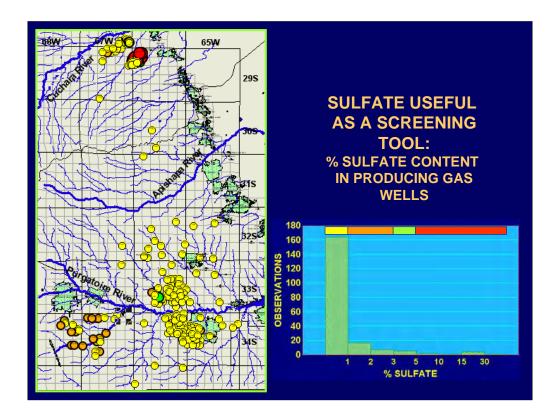
To determine the risk of cross flow between aquifers, we need data that can help us determine the depth of mixing in shallow aquifers. As with most wells in this basin, there is little information available to determine the depth of perforations in a water well. Here, we assume that the perforations are close to the total depth of each well drilled.

Elevated shallow groundwater sulfate concentrations originate from the near-surface oxidation of sulfide-bearing minerals such as pyrite. Such minerals are common in organic-rich shales and in coal seams where they outcrop. Sulfate concentrations decrease with depth principally as the result of sulfate-reducing bacteria. Concentrations also decrease as a result of mixing with discharging intermediate water types of different composition. These data show that mixing of intermediate and shallow aquifers takes place up to a depth of around 400 feet below the surface. Similar mixing depth ranges have been observed by the author in both the San Juan and Powder River Basins.

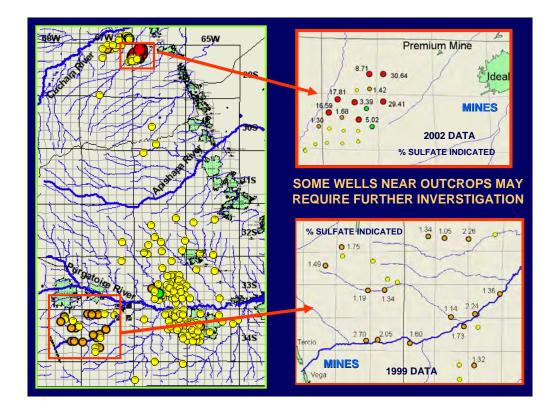
Values for % sulfate shown on the x-axis are calculated as the concentration of sulfate in milliequivalents per liter divided by the sum total of all dissolved anions in milliequivalents per liter ( $Cl^- + HCO_3^- + CO_3^- + SO_4^-$ ). This is the same calculation used to create tri-linear diagrams.



Data published by the USGS in their Water Supply Paper for the basin (#2288) show similar results to those found using Amoco's Spanish Peaks field area data illustrated in the previous slide. Both data sources provide evidence of a floor depth for shallow sulfate on the order of 400 feet below the surface. This would suggest that surface water does not penetrate near-surface aquifers to depths far exceeding 400 feet. The range of 0 - 400 feet is therefore a reasonable estimate for the range across which aquifers can communicate through shallow, permeable fracture networks. Deeper vertical fracture networks are not likely to be as dense nor as permeable as shallow fracture networks because confining stresses increase with increasing depth. We can use such data to estimate that communication between aquifers and deeper producing horizons is not likely to occur in areas where the base of a water well completion is more than 400 feet from the top of a producing well perforation.

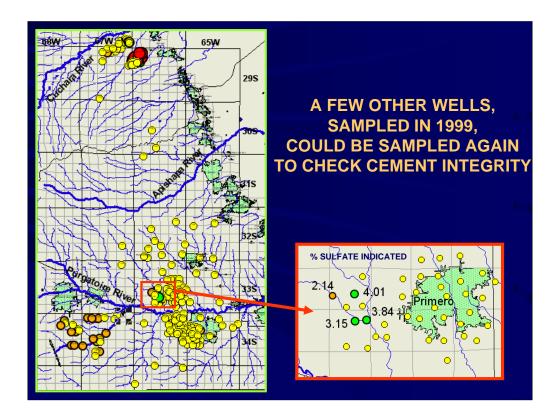


Because most producing wells are completed more than 400 feet below the surface, we can use the relative amount of sulfate in produced fluids as a screening tool to evaluate the risk of communication with shallow aquifers. Mapping relative sulfate concentrations in the Raton basin shows that there are two regions where producing wells contain more than just a trivial amount of sulfate.

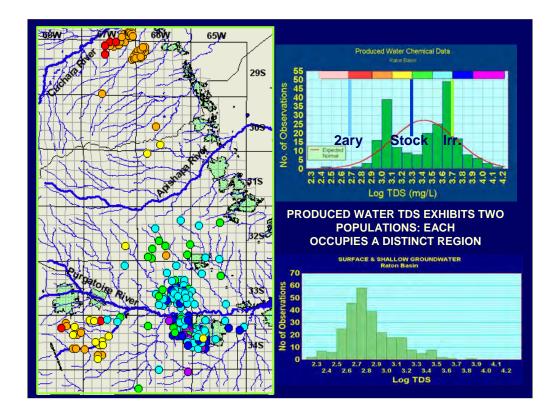


A cluster of northern wells on the eastern side of the basin's synclinal axis have the highest sulfate contents of any produced wells analyzed. These wells produce large volumes of water and relatively little gas. There appears to be a gradient in sulfate contents that increases towards the outcrop belt. This may be an indication that these wells could be in communication with near-surface fluids originating not from overlying aquifers, but from the outcrop belt, or perhaps from the Ideal mine. Alternatively, these wells could be tapping a deeper source of deep-seated sulfate. Further investigation of this area may be necessary to determine if the wells are impacting shallow aquifers along the outcrop. For example, further periodic measurements could show if the relative amounts of sulfate are increasing, or if <sup>13</sup>C DIC values are decreasing.

A cluster of southern wells along the South Fork of the Purgatoire River are on the west side of the basin's synclinal axis. The relative sulfate content in these wells may also be slightly elevated from recharging groundwaters to the west. Elevated sulfate concentrations in this area are not nearly as high as those observed in northern wells. Produced fluids in this area were sampled in 1999, and could be sampled again to determine if there are systematic changes in relative sulfate content. If water production volumes have not declined significantly, this too could be an indication of communication with other aquifers.

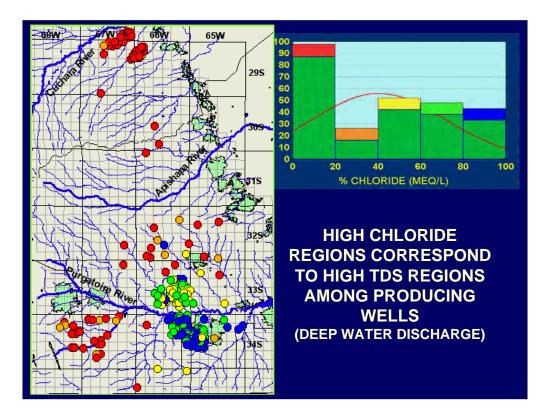


KLT wells such as the Tandy 13-21, Susan 33-21, Jenni 33 - 28, and the Lucy 41-28 show elevated relative concentrations of sulfate when last sampled in 1999. If these wells still produce water, these wells could be sampled again to determine if the percent sulfate has either increased or decreased since 1999. If sulfate levels have increased or remain steady, it might be useful to check the wells for cement integrity. Water production data may provide useful indicators of possible communication with other aquifers if high volumes persist over time.

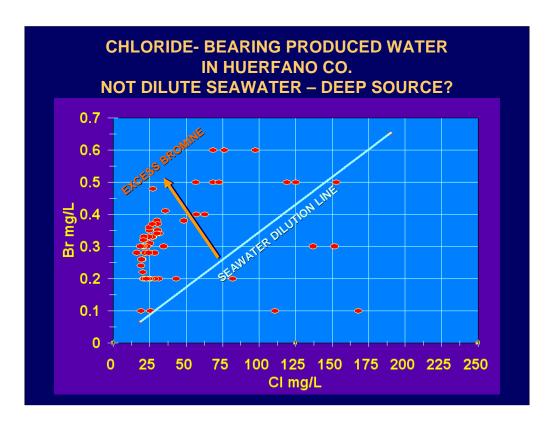


TDS values are usually reported after measuring dry weights on filter paper once a fixed amount of fluid is allowed to evaporate. Loss of CO<sub>2</sub> during this process results in TDS values being lower than the sum of dissolved constituents analyzed using ion chromatography (IC), atomic adsorption (AA), and inductively coupled plasma (ICP) techniques. This method is appropriate when attempting to evaluate the environmental impact of discharging produced fluids to the surface. For purposes of evaluating *in-situ* values, we have chosen to sum up all the values reported for the major dissolved ions in mg/L. This convention provides more consistent values as a basis for comparing samples. Analysis of available data shows that there tends to be a 5% to 10% discrepancy between TDS values measured by evaporation, and TDS values computed on the basis of summing dissolved ion contents.

TDS values in the basin are variable, and exhibit a bimodal distribution. The second mode of high TDS values are concentrated along the Purgatoire River drainage basin, just west of the Trinidad Reservoir. Coalbed gas wells in this region produce the lowest water rates, and the highest gas rates.



Values for % chloride are calculated on the same basis as that used to create tri-linear diagrams. The value represents the concentration of chloride in milliequivalents per liter divided by the total milliequivalents of anions ( $Cl^- + HCO_3^- + CO_3^- + SO_4^-$ ). High chloride regions correspond to regions of high total salinity. The marked difference in chemical composition between areas of high salinity (high chloride content) and low salinity (low bicarbonate content) suggests that the south-central area of the basin has not been flushed by recharging, relatively low salinity bicarbonate waters. This area may be a discharge zone for deep circulating brines or, alternatively, an area in hydrologic isolation from the rest of the groundwater fluids in the basin.

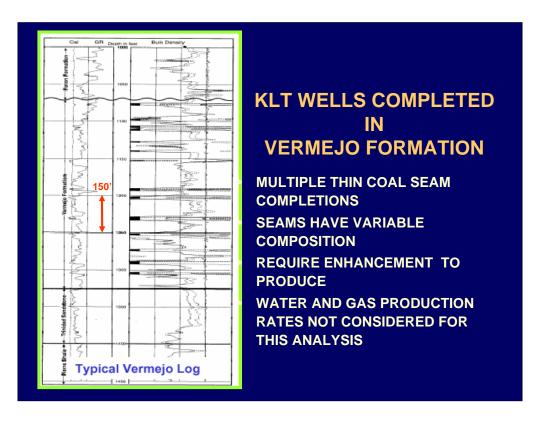


There are only a limited amount of data available recording the bromine content in produced water samples. Most are from wells in Huerfano County. These can be used to determine the origin of fluids. If chloride and bromine are derived from diluted seawater, as may be present in marine sediments within and below the Trinidad Sandstone Formation, then the ratio of bromine to chlorine should be constant. Constant ratios would fall along a seawater dilution line as indicated. However, most samples exhibit an excess bromine concentrations above that which would be present if the fluids contained diluted sea water. Although such a chemical signature is usually characteristic of fluids which have dissolved evaporite salt deposits, there is no published evidence to indicate the presence of such deposits in the sediments of this basin. It may be that the source of bromine is somehow related to the emplacement of the Spanish Peaks intrusive complex. Whatever the source, excess bromine is an extraneous component that is indicative of deeply circulating brines.

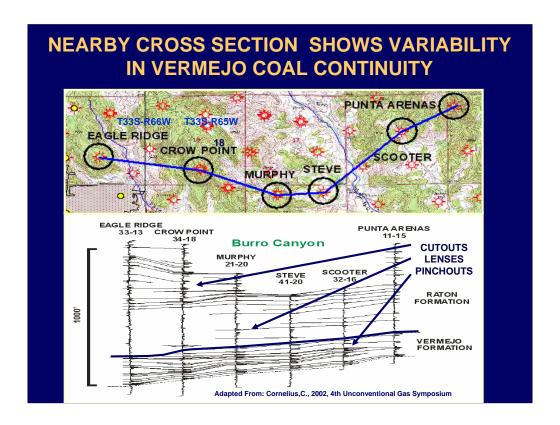
# SPATIAL AND TEMPORAL STUDY OF KLT PRODUCING WELLS: VISUALIZING AQUIFER RESPONSE

- 95 PAIRS OF SAMPLES EVALUATED
- AVERAGE LAG TIME BETWEEN SAMPLES 3 YRS. FROM 1999 - 2002
- EARLY DATA SET CONTAINS ALL MAJOR IONS
- LATE DATA SET CONTAINS LIMITED DATA INCLUDING DISSOLVED CHLORIDE & BICARBONATE CONTENT
- DATA USEFUL TO EXAMINE
  - COMMUNICATION WITH SURFACE
  - PRODUCTION COMPARTMENTATION
  - REGIONAL FLOW RATE BOUNDARIES
  - LOCAL FLOW RATE BOUNDARIES

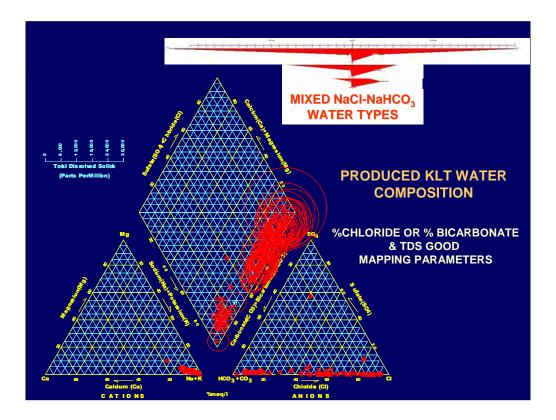
The principal objective of this study is to determine if produced water chemistry can help establish the spatial range of connectivity among the gas productive water-bearing horizons. Methods used to assess reservoir continuity using geochemical data rely on examining both spatial and temporal trends.



Coal seams in the Vermejo Formation are largely thin and discontinuous. Published geologic studies in the area commonly note great difficulty in correlating seams beyond distances of one to two miles. There are also other factors that isolate the production behavior of producing coalbed gas wells. Such factors include: variable formation damage due to drilling, variable effectiveness of completions in each seam, structural discontinuities along minor faults that generally offset coals across vertical distances between 25 to 50 feet, and discontinuities introduced by volcanic intrusions such as dikes and sills. Published data show that each coal seam interval must be hydraulically stimulated to produce effectively. Tracer analyses further show that induced hydraulic fractures do not extend far beyond the vertical intervals perforated. Such data provide important evidence documenting the limited extend of vertical communication between coal seam aquifers.



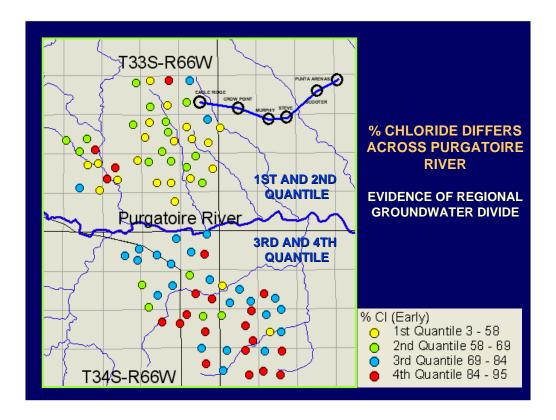
This cross-section, just to the east of the KLT wells to be illustrated next, identifies the stratigraphic variables governing the heterogeneity of coal seams and other sediments in the area. Each square in the diagram represents a section. These discontinuous coals seams are encased in impermeable shale layers and cemented sandstone layers.



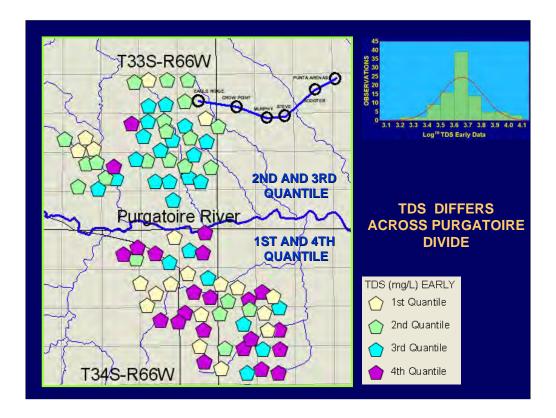
Early data collected by KLT and analyzed sometime in 1999 contain complete analyses of the major ions. Results of these analyses are illustrated above. Late data collected in 2001 and 2002 only contain quantitative analyses of chlorides, bicarbonates, and total dissolved solids.

The trend observed among KLT producing wells is the same as the trend observed for all samples of produced water. Formation fluids are principally composed of mixtures of sodium bicarbonate and sodium chloride. Accordingly, by plotting either percent chloride or percent bicarbonate, one can fully represent the chemical composition of produced fluids.

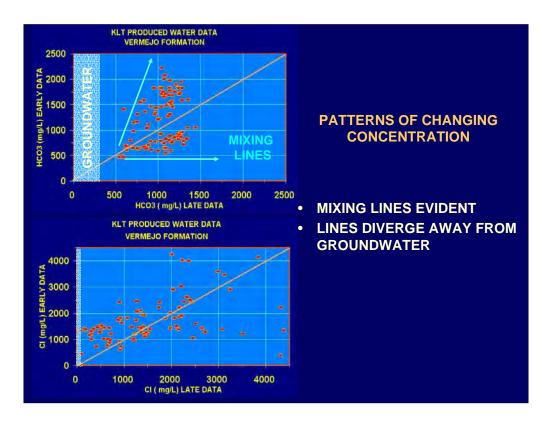
Two types of compositional plots can help us understand aquifer compartmentation in the basin. The spatial distribution of values obtained early in the production history of a field is best used to observe the net result of long term mixing between aquifers. Temporal chemical changes measured over the short term during production are better suited to show both the spatial distribution of local flow units and the effectiveness of well bore completions.



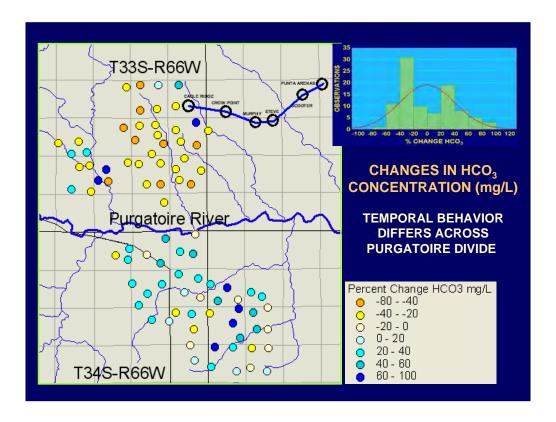
The spatial distribution of produced water chemistry evolved over geologic time is illustrated above. To simplify results, % chloride has been color coded to represent quartiles in the relative chloride content as measured among producing KLT wells in 1999. It is evident that the Purgatoire River drainage divide separates groundwaters into two distinct hydrologic regimes. Chloride values in the 1<sup>st</sup> and 2<sup>nd</sup> quartiles dominate the northern half of the divide, whereas chloride values in the 3<sup>rd</sup> and 4<sup>th</sup> quartiles dominate the southern half of the divide. There is also a fair amount of spatial discontinuity among chloride values on each side of the divide. Such data provide evidence of both regional and local aquifer discontinuities.



TDS values collected in 1999 have also been simplified in terms of quartiles in the above illustration. The trends observed in TDS are distinctly different than the trends observed using normalized % chloride values. North of the Purgatoire River divide, TDS is dominated by values in the 2<sup>nd</sup> and 3<sup>rd</sup> quartiles; south of the divide, TDS is dominated by values in the 1<sup>st</sup> and 4<sup>th</sup> quartiles. A large amount of local heterogeneity is also evident.

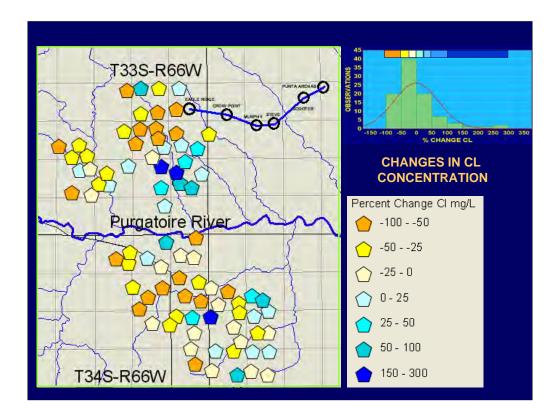


Temporal production trends showing differences in produced water chemistry between 1999 and 2002 are illustrated above. The patterns of changes in bicarbonate and chloride concentrations are substantially different. In the upper plot of bicarbonate values, we can observe that the data split into two groups. The first group of wells, above the orange reference line of no change, show an increase in bicarbonate concentration with time. The second group, plotting below the orange line, shows a decrease in bicarbonate concentration with time. The change in chloride values illustrated in the lower plot do not show the bimodal behavior evident in bicarbonate values. This indicates that the factors governing bicarbonate concentrations are independent from those governing chloride values. This would be expected because the origins of these fluids vary. It is possible that changes in bicarbonate concentrations mainly reflect the variable input of water produced from coal seams, whereas variable chloride concentrations mainly reflect the variable input of water produced from permeable sandstones. The basis for this statement merely relies on the observation that coal seams all produce variable amounts of CO<sub>2</sub> gas, a gaseous component that dissolves in water to form bicarbonate. Both diagrams show that mixing fluids have a trend which is opposite in direction to the trend that would be observed if fluids were mixing with shallow groundwaters.



Percent bicarbonate values used for this map are derived by subtracting the latest measured bicarbonate concentrations (in years 2001 and 2002) from the earliest measured concentrations (3 in 1989 and the rest during 1999), and dividing by the early values ((mg/L HCO<sub>3</sub>- yr. 2002 - mg/L HCO<sub>3</sub>- yr. 1999)/(mg/L HCO<sub>3</sub>- yr. 1999))\*100. This approach normalizes the changes in bicarbonate composition during an average production interval of 3 years.

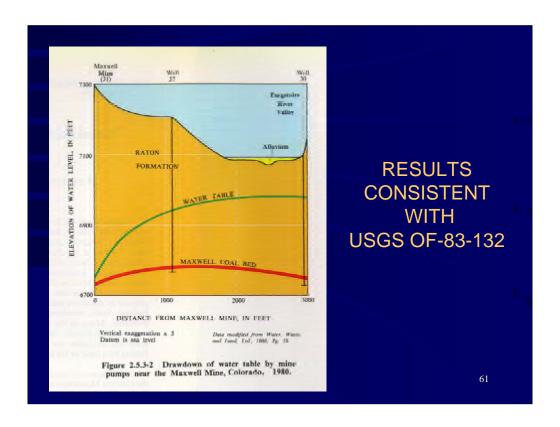
Results show that produced fluids north of the Purgatoire River respond differently to short term production activities than produced fluids south of the river. This confirms that the river is a groundwater divide where fluids recharged from the northeast meet fluids of differing composition that are recharged from the southwest. It is also possible that there may be measurable differences in the thickness and distribution of Vermejo Formation coal seams on either side of the river. Small-scale local differences in produced fluid response is also evident, indicating that changes in produced water chemistry cannot be correlated beyond distances of between one to 1.5 miles.



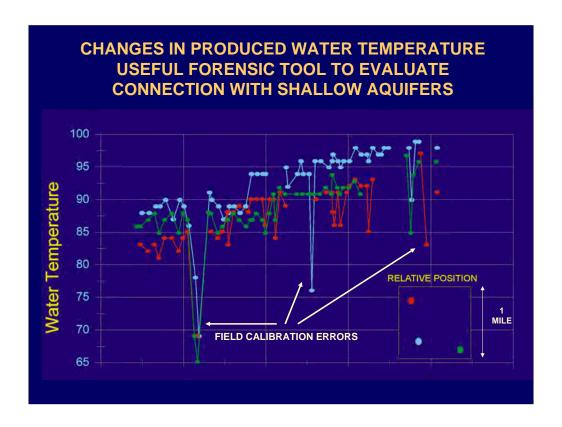
The percent change in chloride concentrations was calculated by subtracting early measurements from late measurements (in mg/L) and dividing them by the early measurements. As noted earlier, changes in chloride concentrations do not exhibit the same regional and local patterns as changes in bicarbonate concentrations. This supports the concept that chloride-bearing fluids may originate from different sources, such as permeable sandstones.

Maps such as these could be combined with water production data for a variety of uses. For example, they may be used to examine the effectiveness of completions in coal. For example, areas where the bicarbonate concentration decreases while the chloride concentration increases may be indicative of sandstones that are contributing significantly to water production. The highly heterogeneous response also has implications to those attempting to explore for gas accumulations in sandstone bodies. These data indicate that values of formation water resistivity are likely to be very different among individual sandstone bodies. Accurate assessment of formation water resistivity is of paramount importance when attempting to predict gas saturation in sandstones.

The main conclusion to be drawn from mapping spatial and temporal maps of dissolved constituents in produced water is that visible patterns are clustered. Clusters show a different degree of spatial coherence at both the regional and local level. At the local level, it is not possible to correlate produced water patterns beyond distances of between one to one and a half miles. We can reasonably conclude from these data that the aquifer continuity in this portion of the Raton Basin is limited, and that production practices are not likely to influence drawdown over an area larger than 1 to 1.5 miles. Although results of pressure interference testing could add more information to the degree of communication between adjacent wells, such data are not available in the public domain.



Data published by the USGS show that water levels in the Maxwell coal seam, declining as a result of mine pumping, can be affected to distances approaching ½ a mile from the mine. Such corroborating evidence confirms the relatively limited extent of aquifer continuity in the basin.



Water composition data are not the only types of measurements useful for examining fluid flow in producing coalbed aquifers. These monthly temperature measurements, obtained from producing wells in Huerfano County, were made using a thermistor device. Unfortunately, the thermistor was not regularly calibrated. Calibration errors are evident as strong excursions from the observed long term trends. The relative position of wells in a one square mile section is indicated in the reference box at the lower right hand side of the diagram. Regular measurements of produced water temperatures can be most useful for determining potential communication with the surface. Long term cooling trends would indicate possible communication with relatively cool near-surface waters. As evident here, the long term trends indicate admixture with warmer waters. Because topographic relief and therefore formation temperature increases to the west, this diagram would indicate a net flow of fluids from the west. Excursions from long term trends measured in real time should provide an important screening tool necessary to sample fluids for complete chemical analysis.

## SUMMARY: RELEVANT FEATURES OF INTERMEDIATE AND DEEP GROUNDWATER SYSTEM

- PURGATOIRE RIVER IS A GROUNDWATER DIVIDE
- GEOLOGIC STRUCTURE AND TOPOGRAPHY TOGETHER ALLOW DISCHARGING DEEP AND INTERMEDIATE GROUNDWATERS TO MIX WITH RECHARGING SURFACE WATER
- MIXING RATES AND ROCK-FLUID INTERACTION RATES VARY WITH GEOLOGY AND ARE SPATIALLY VARIABLE
- STRATIGRAPHY AND LOCAL DEFORMATION IS COMPLEX LARGE REGIONAL AND LOCAL FLUID DISPERSION EXPECTED
- WATER FROM PRODUCING ZONES POORLY MIXED, INDICATING POOR VERTICAL AND LATERAL CONTINUITY
  - CONSISTENT WITH PRODUCTION PRACTICES
  - CONSISTENT WITH CONCLUSIONS IN USGS WSP 2288
- RISK OF GAS OPERATIONS TO SHALLOW GROUNDWATER RESOURCES LOW
  - WHERE VERTICAL SEPARATION BETWEEN INTERVALS >400'
  - WHERE LATERAL SEPARATION > 1.5 MILES

#### **SUMMARY:**

### USEFUL SCREENING TOOLS TO ADDRESS POTENTIAL IMPACT ON SHALLOW AQUIFERS

- TRILINEAR DIAGRAMS VIEW MIXING TRENDS AND SELECT MAP PARAMETERS
- SPATIAL MAPS LARGE SCALE HETEROGENEITY
- TEMPORAL MAPS SMALL SCALE HETEROGENEITY & CONNECTIVITY OF AQUIFERS
- SULFATE CONTENT: SURFACE WATER MARKER
- 13C DIC- HCO<sub>3</sub>: SURFACE MARKER, SOURCE MARKER
  - (SPATIAL & TEMPORAL)
- PRODUCED WATER TEMPERATURE
- COMPLETION INFORMATION
  - BASE WATER WELL PERFORATION
  - TOP PRODUCING WELL PERFORATION

#### **CONCLUSIONS**

- RISK OF POTENTIAL IMPACT TO SHALLOW AQUIFERS CAN BE SCREENED AS FOLLOWS:
  - MONITOR SHALLOW AQUIFER WATER LEVELS WHERE NEEDED
  - MAP LATERAL AND VERTICAL PROXIMITY OF GROUNDWATER AND PRODUCED WATER COMPLETIONS
  - ANALYZE PRODUCED WATER PROPERTIES
    - REGULAR, CALIBRATED TEMPERATURE SURVEYS
    - INITIAL MAJOR ION ANALYSES
      - SUPPLEMENTED BY ADDITIONAL TYPES OF ANALYSES AS NEEDED BASED ON ACTION TRIGGERS
    - PERIODIC MONITORING OF CHANGES IN CHEMICAL PROPERTIES
    - ANALYZE STABLE ISOTOPIC SIGNATURE OF DISSOLVED BICARBONATE
    - MONITOR CHANGES IN TDS AND SULFATE CONTENT
- RISK OF SEEPS CAN BE SCREENED WITH IRS AND DISSOLVED METHANE MEASUREMENTS
- THE ORIGIN OF METHANE CAN BE DETERMINED USING STABLE ISOTOPE MEASUREMENTS